Variabilidad atmosférica y ciclones tropicales



Atlantic SST Multidecadal Mode 1870-2004



Tropical North Atlantic SSTs Annual 1856-2004



Huracanes en el Caribe: Oscilación Multi-decadal del Atlántico



August-October Climatology 200-850 hPa Vertical Shear August-October Climatology 200 hPa Streamlines Zonal Wind





(Landsea, 2005)

EQ

120W

100W

8ÓW

60W

4Ô₩

2ÓW

La Niña versus El Niño: Probabilidad de formación de huracanes



Florida, then became a depression in the Gulf of Mexico before regaining TS strength and finally making landfall as a HR. (i.e., there are really only 4 storms on the plot.

Cizalla del viento: Niño versus no - Niño

Variabilidad en el Pacifico Oriental: número de ciclones

Table I. Statistics of cyclone activity from 1966 to 2004 in the NE Pacific. The numbers in parentheses represent the mean number of storms per year.

	Tropical storms	Hurricanes					
		Cat. 1	Cat. 2	Cat. 3	Cat. 4	Cat. 5	Total
El Niño years	98	45	19	27	37	5	231
	(6.53)	(3.0)	(1.3)	(1.8)	(2.5)	(0.33)	(15.4)
Non-El Niño years	182	88	40	33	41	4	388
	(7.9)	(3.8)	(1.7)	(1.4)	(1.8)	(0.17)	(16.9)
Total	280	133	59	60	78	9	618
	(7.4)	(3.5)	(1.5)	(1.6)	(2)	(0.23)	(16.3)

(Romero-Vadillo et al, 2007)

Variabilidad en el Pacifico Oriental : duración de los ciclones

Table II. Statistics of cyclone duration during the period 1966 to 2004 in the NE Pacific. The numbers in parentheses represent the mean number of storms in each category.

		Days				
	1 to 3	4 to 6	7 to 9	10 to 12	Longer	Total
					than 12	
El Niño	53	92	59	31	14	248
years	(3.5)	(6.1)	(3.9)	(2.1)	(0.9)	(16.5)
Non-El Niño	85	143	96	40	7	371
years	(3.7)	(6.1)	(3.9)	(2.1)	(0.9)	(16.13)
Total	138	234	155	71	21	619
	(3.6)	(6.2)	(4.1)	(1.9)	(0.6)	(16.3)

(Romero-Vadillo et al, 2007)

Variabilidad en el Pacifico Oriental

(Romero-Vadillo et al, 2007)

The "Global Mode": Enhanced 200mb Subtropical Ridges

Active minus Inactive: August - September 200 mb Streamfunction, Winds

200-hPa Streamfunction Anomaly

"Global Mode": Winter '98-'99 To '99 Hurricane Season

"Global Mode": Winter '96-'97 To '97 Hurricane Season

A continuación se incluyen dos presentaciones de colegas en un congreso reciente sobre ciclones tropicales.

1st International Summit on Hurricanes and Climate Change (27 May – 1 June 2007, Crete, Greece) 1st International Summit on Hurricanes and Climate Change (27 May – 1 June 2007, Crete, Greece)

Global warming and tropical cyclone climatology as simulated in a 20-km-mesh global atmospheric model

J. Yoshimura (MRI) K. Oouchi (AESTO/MRI, FRCGC) H. Yoshimura (MRI) R. Mizuta (AESTO/MRI) S. Kusunoki (MRI) and

20-km-mesh atmospheric GCM simulation

- Unprecedented as a global climate simulation.
- The main focus is TC climatology under global warming condition in the end of 21st century.
- SST given as a boundary condition:
 Excludes coupling with ocean subsurface,
 Enables much-higher-resolution simulation.

Early results of this study have been published.

"Earth Simulator" – A huge computing system in Yokohama, Japan –

Super-high-resolution global atmospheric model

- Collaborative work
 - NPD/JMA
 - Climate Research Department of MRI
- JMA : Next operational global NWP model
- MRI : Next generation climate model
- Based on operational JMA-GSM
- Fortran90 coding style
 - Resolution:TL959(20km) with 60 layers
 - Time integration: Semi-Lagrangian Scheme (Yoshimura, 2004)
 - SW radiation: Shibata & Uchiyama (1992)
 - LW radiation:Shibata & Aoki (1989)
 - Cumulus convection: Prognostic Arakawa-Schubert (Randall and Pan, 1993)
 - Land hydrology: Sib with 4 soil-layers and 3 snow-layers
 - Clouds: large-scale condensation, Cumulus, stratocumulus
 - PBL: Mellor & Yamada (1974,1982) level-2 closure model
 - Gravity wave drag: Iwasaki et al. (1989) + Rayleigh friction

The same model used for NWP

(09 Jul 2002 00UTC, FT=24, GMS-IR forecast images)

The same model used for NWP

"Time-slice" simulations of future climate change based on the IPCC SRES A1B scenario.

SST data from two coupled AOGCMs

- MRI-CGCM2.3.2
 - AOGCM of MRI
 - Resolution: atmosphere T42L30, ocean 2.5°×
 2°
 - With "flux adjustment"

- **MIROC** ver 3.2 high-resolution
 - AOGCM of CCSR, NIES and FRCGC
 - Resolution: atmosphere T106L56, ocean

SST settings

"Present-day" End of 20 th	"F uture " End of 21 st century		
century Obs. SST	Obs. SST MRIASST 20 yrs	SST	
20 yrs	Obs. SST MIROC ASST 10 yrs	anomalies	
MRI SST 20 yrs with year-to-year variation	MRI SST 20 yrs with year-to-year variation	"Raw"	
MIROC SST 5 yrs	MIROC SST 5 yrs	SSTs	

Seasonal-mean precipitation

Model: TL959L60

OBS: CMAP

Latitude

16

8

2 ۵

Criteria to define TCs in the model

Based on those of Sugi et al. (2002) and Bengtsson et al. (1996)

- [1] A grid point (45N-45S) with local-minimum surface pressure is selected as a TC-center candidate. The minimum surface pressure is at least 2 hPa lower than mean surface pressure of a 7°×7° grid box.
- [2] Maximum magnitude of relative vorticity at 850 hPa is above 3.5x10⁻⁵ s⁻¹ near the point.
- [3] Maximum wind speed at 850 hPa is larger than 15 m s⁻¹ near the point.
- [4] Temperature structure has a warm core above the point: Sum of temperature deviations at 300, 500 and 700 hPa exceeds 2.0 °C
- [5] Maximum wind speed at 850 hPa near the point is larger than that at 300 hPa. (This criterion is to remove extratropical cyclones.)

Tracks of tropical cyclones (TCs)

TC frequency as a function of peak wind intensity

Global-mean SST warming MRI \triangle SST = 1.6 vs. MIROC \triangle SST = 3.2

Increase of intense cases under global warming

TC frequency as a function of peak wind intensity

The number of TC formation

The number of TC formation

W.N. Pacific: the number of TC formation (0N 45N, 100E 180E # per year

	"Present-day"	"Future"	
	End of 20 th	End of 21 st century	Change
TZZ	Eobs. SST]	Eobs-SST MRIASST] 8-1	-4.5
s used	7 5 -6	EObs. SST MIROC ASST] 15.9	+3.3
"Raw" data of SST used	EMRI SSTJ 20-9	EMRI SST]]].4.9	-6-0
	EMIROC SSTI 15-8	EMIROC SSTJ 5-8	-10-0

W. N. Pacific: climatological precipitation (mm/day) Jul-Aug-Sep

Present: Obs. SST expt.

1 2 4 6 8 10 15 20

Future: Obs. SST MIROCASST

Summary

- Frequency of tropical cyclone formation decreased globally by ~30%.
- Frequency of <u>intense</u> tropical cyclones increased with a higher SST condition.
- Basin-scale tropical cyclone activity <u>depends</u> on regional SST variations.
 - Consistency of a constructed SST condition with its original AOGCM is discussed.
 Such consideration could enable more robust projection of regional TC activity.

RESPONSE OF TROPICAL CYCLOGENESIS TO GLOBAL WARMING IN IPCC AR-4 SCENARIOS ASSESSED BY A MODIFIED YEARLY GENESIS PARAMETER

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The Yearly Genesis Parameter of W. Gray

The Yearly Genesis Parameter (YGP) represents cyclogenesis in an oceanic basin (Gray, 1975) as the sum of 4 the Seasonal Genesis Parameters (SGP) that are the product of a thermal and dynamical potential.

Thermal potential :	<pre>{ ocean heat content vertical stability of the lower troposphere mid-tropospheric moisture</pre>
Dynamical potential :	Coriolis effect low layer vorticity vertical shear of the horizontal wind

Simulated YGP and CYGP as compared to observations (Royer et al., 1998)

Figure 8. Geographical maps of the convective YGP calculated from ARPEGE-Climat control experiment CO1 (units: number of TC genesis in 20 years per 5° lat.-lon.). Isolines are drawn at 1, 5, 10, 20, 40, 50 and values above 5 are shaded.

Figure 3. Geographical map of the YGP (a) calculated from ARPEGE-Climat control experiment (CO1), (b) as observed from 1973 to 1992 (Tsutsui and Kasahara, 1996) (units: number of TC genesis in 20 years per 5° lat.-lon.). Isolines are drawn at 1, 5, 10, 20, 40, 50 and values above 5 are shaded.

Components of the dynamical potential (JAS season)

Shear:

 $(200 - 850 \text{ hPa horiz. wind shear in ms}^{-1}/650 \text{ hPa})$

$$I_{s} = 1/(|dv/dp| + 3)$$

Vorticity: (850 hpa rel. vort ζ in 10⁻⁶ s⁻¹) $I_v = max(0, \zeta (f/|f|) + 5)$

Dynamical potential: ($f=2\Omega \sin \phi$ in 10^{-5} s^{-1})

$$\mathsf{DP} = \mathsf{I}_{\mathsf{s}} \cdot \mathsf{I}_{\mathsf{v}} \cdot \mathsf{f}$$

Dynamical and convective components of the CYGP (JAS season)

Dynamical potential:

$$\mathsf{DP} = \mathsf{I}_{\mathsf{s}} \cdot \mathsf{I}_{\mathsf{v}} \cdot \mathsf{f}$$

Convective potential: (conv. Precip. P_c above threshold P_0)

 $CP = k. max(0, P_c - P_0)$ (k calibration factor -> 84 TC genesis annually on average)

Genesis potential: SGP = DP. CP

Model CYGP 1970-2000

